THE EXTENSION OF THE CALAVERAS - SHOO FLY THRUST TO THE SOUTHERN END OF THE SIERRA NEVADA METAMORPHIC BELT, CALIFORNIA

Preliminary geologic maps of the Tuolumne, Duckwall Mountain, Groveland, Jawbone Ridge, Lake Eleanor NW, and NE 1/4s, Buckhorn Peak, Kinsley, El Portal, and Buckingham Mountain 7 1/2 minute quadrangles, California Division of Mines and Geology, Sacramento, California. Accompanied by 10 p. report plus 5 figures, 1 table, unpublished DMG report.

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ABSTRACT

THE EXTENSION OF THE CALAVERAS-SHOO FLY THRUST (CSFT) TO THE SOUTHERNEND OF THE SIERRA NEVADA METAMORPHIC BELT (SNMB), CALIFORNIA MERGUERIAN, Charles, Hofstra University, Geology Department, Hempstead, NY, 11549 and Lamont Doherty Geological Observatory of Columbia University, Palisades, NY, 10964.

The CSFT has now been mapped within the SNMB as far south as Lat.37°30'N in Mariposa County, where the belt is engulfed by the Sierra Nevada batholith (SNB). The thrust has been folded and overturned north of the Tuolumne R. by forceful intrusion of the 170 m.y. Standard pluton. It is truncated by granite (SNB) south of Hunter Bend of the Clavey River and re-emerges trending parallel to, but northeast of Pilot Ridge. The CSFT is cut by the Hazel Green pluton and then, with a more southerly trend, extends 16 km beyond the Merced River. The sub-vertical ductile thrust separates polydeformed quartzite, mica-quartz schist \pm garnet and graphite, calc-silicate rock, and augen gneiss of the upper plate Lower Paleozoic Shoo Fly Complex on the east from argillite, chert, marble, and talc-schist of the lower plate Calaveras Complex on the west. Within the 250 m-wide CSFT zone, syn-metamorphic imbrication of rock units and intense silicification is common. Slices of the Calaveras occur as flattened and disarticulated meter-scale sheets of recrystallized argillite, massive and rhythmic-bedded chert, quartz sandstone and mylonitic marble intruded locally by syn-tectonic granite. In the Shoo Fly, the CSFT fabric (S_3) is marked by the development of blastomylonite with 0.1-1.0 cm-spaced syn-metamorphic shears and attendant isoclinal folding and transposition of pre-thrust penetrative fabrics (S₁ and S₂) into S₃. Since the Shoo Fly, with its distinctive structural style, occurs in the upper plate of the CSFT as far south as Lat. 37°30'N, roof pendants to the east and southeast must be restudied to ascertain the full extent of Lower Paleozoic wallrocks and the tectonic significance of the CSFT.

22 December 1981

INTRODUCTION

The Sierra Nevada range (Figs. 1, 2) has long presented almost insurmountable obstacles to an understanding of the Paleozoic and early Mesozoic tectonic history of the southwestern margin of the Cordillera, largely be-cause the massive Sierra Nevada batholith has obliterated a major part of the evidence and separates sparsely fossiliferous rock sequences in the western and eastern parts of the range. Most previous studies have either been concentrated in small areas where local geologic relations could be established, or have been reconnaissance studies lacking-in detail.

In Tuolumne County, California, recent work (Fig. 3) has established a stratigraphic and structural data base for comparison to other regions and has shown the presence of two subparallel belts of rock (the Calaveras and Shoo Fly Complexes) and the presence of a ductile thrust between them.

Detailed and reconnaissance bedrock mapping at a scale of 1:24,000 was conducted in the region southeast of the Standard pluton in July and August, 1981. The work was concentrated on the San Jose and Mariposa Sheets. Figure 4 shows the USGS quadrangles that were mapped and included with this report. The quadrangles mapped were: Tuolumne, Duckwall Mountain, Groveland, Jawbone Ridge, Lake Eleanor (NW, SW, SE), Kinsley, El Portal, and Buckingham Mountain The last three 7 1/2' quadrangles have been drafted onto a 15'base map (El Portal 15" quadrangle). After some general remarks on the geology of the Shoo Fly Complex I will discuss the geologic discoveries generated by this summer's research.

PREVIOUS WORK IN THE SHOO FLY

The Shoo Fly contains at least two generations of structures that pre- date the development of the Calaveras - Shoo Fly Thrust (Merguerian and. Schweickert, 1980). Earliest recognizable are a metamorphic layering (S_1) and isoclinal folds of bedding that are crosscut by deformed granitoids - augen gneiss. A second generation of structures (S_2) created a penetrative foliation in the augen gneiss, isoclinal folds of layering and S_1 in the quartzites, and a strong mylonitic fabric. Both early sets of structures are deformed by third-generation structures (S_3) that are related temporally to the Calaveras Shoo Fly Thrust.

Figure 3 is a generalized geologic map of the Shoo Fly in northern Tuolumne County, California showing the widespread occurrence of augen gneiss. Preliminary U-Pb zircon analysis by Jason Saleeby (personal communication) indicates that the augen gneisses may have igneous ages of 400 m.y. and older, and that the gneisses have undergone a major Triassic or older metamorphic event. If true, the first generation structures above are early Paleozoic and the second and/or third structures could be Triassic or older. The structural chronology for the Shoo Fly and Calaveras is shown in Table 1 and is explained in the two-page legend accompanying the maps. Figure 5 is a cartoon that summarizes the structural relations.

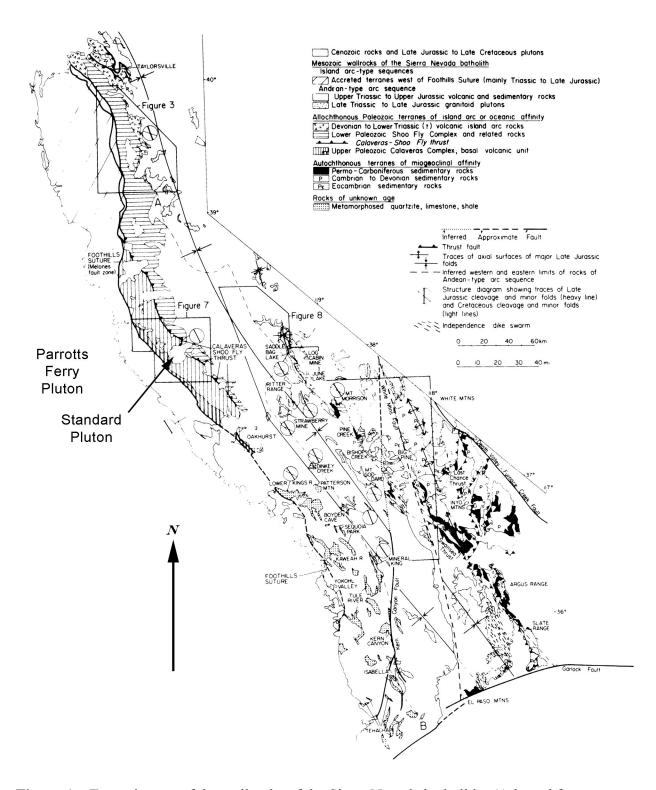


Figure 1 – Tectonic map of the wallrocks of the Sierra Nevada batholith. (Adapted from Schweickert, 1981, Fig 5-1, p. 91,)

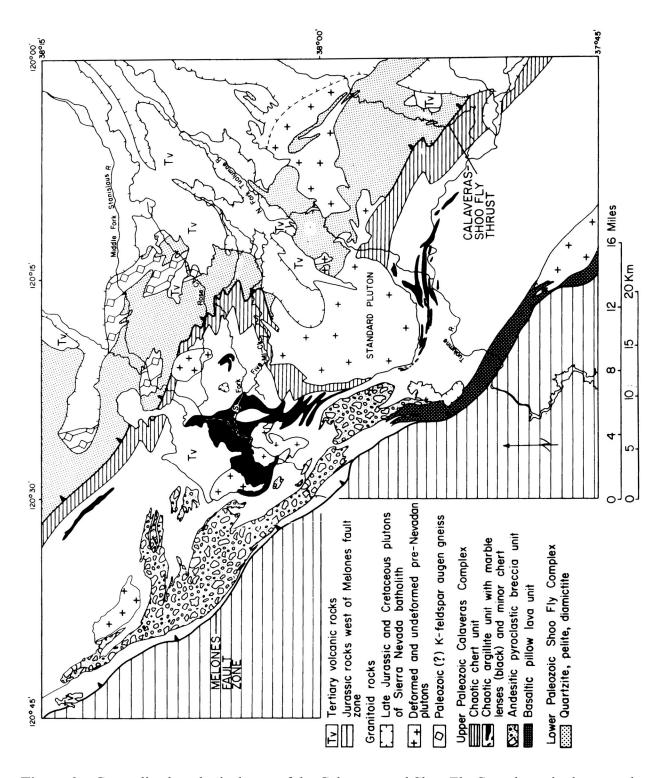
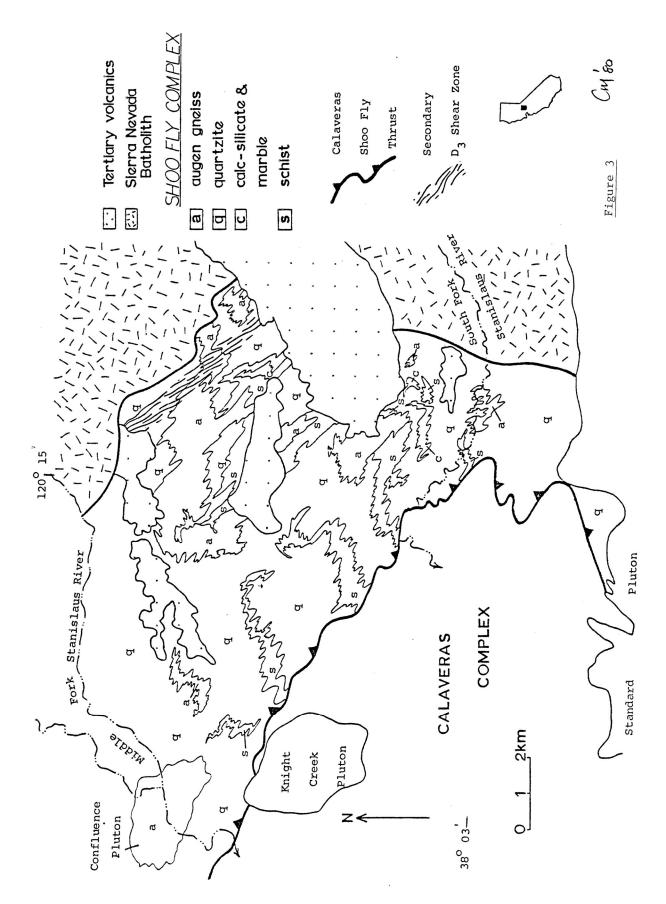
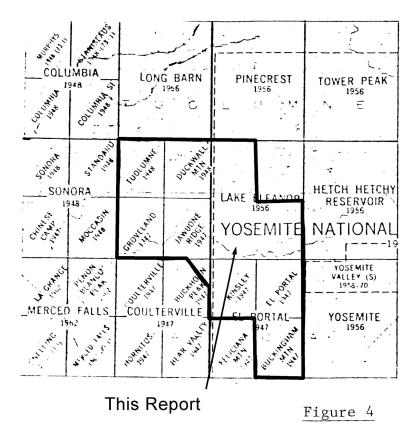


Figure 2 – Generalized geological map of the Calaveras and Shoo Fly Complexes in the central Sierra Nevada, California. (Adapted from Schweickert, 1981, Fig 5-7, p. 104.)





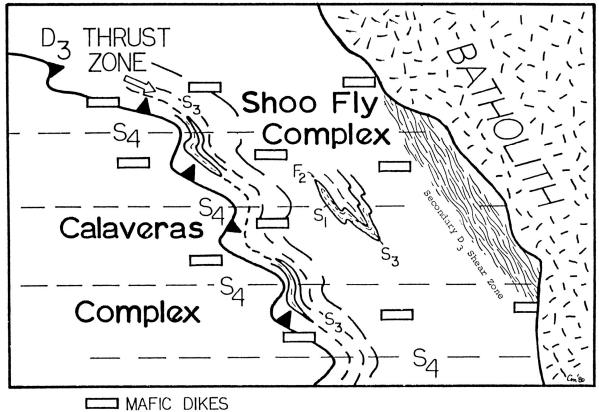


Figure 5 - Cartoon showing structural relations of basement rocks in Tuolumne Co., California

THE CALAVERAS-SHOO FLY THRUST

The Calaveras-Shoo Fly Thrust is an east-dipping mylonitic thrust zone which separates two distinctive Paleozoic bedrock complexes over its known extent of 175 km in the western Sierra Nevada foothills of California (Fig. 2). The upper plate rocks of the lower Paleozoic Shoo Fly Complex record a longer and more complex deformational history than the lower plate rocks of the upper Paleozoic (?) Calaveras Complex. In particular, the Shoo Fly exhibits two penetrative structural events that formed prior to juxtaposition with the Calaveras (Meguerian, 1981). The thrust formed between early Permian and middle Jurassic time, probably during the Permo-Triassic Sonoma orogeny and represents a deep-seated dislocation between a late Paleozoic island arc massif and its associated subduction complex (Schweickert and Merguerian, in press). A more detailed discussion of the structural history of the Shoo Fly is included in the 1981 RFP Proposal. The key to unraveling the complex geologic history of the southern Sierra Nevada foothills was accomplished by tracing the relevant structural and stratigraphic data southward from Tuolumne County, in accord with my 1981 proposal and subsequent contract. The results of this research is presented, together with maps, below.

Table 1 - Tentative structural chronology Calaveras and Shoo Fly complexes.

Calaveras Complex	Ages	Shoo Fly Complex
S ₅ spaced cleavage (N70°W)	Cretaceous?	S ₇ crenulation cleavage (N70°W)
S ₄ crenulation cleavage (N20°E)	Late Jurassic?	S ₆ crenulation cleavage (N20°E)
S ₃ crenulation cleavage	Late Jurassic	S ₅ crenulation cleavage
(N20°-30°W)	(Nevadan Orogeny?)	(N20°-30°W)
	++++ Intrusion of Sonora	
	mafic dike swarm ++++	
	++++ Intrusion of Parrotts	
	Ferry pluton (170 m.y. ago)	
	++++	
S ₂ spaced cleavage and mica schistosity (N70°W)	Early Mesozoic	S ₄ spaced cleavage and mica schistosity (N70°W)
S ₁ flattening foliation and	Permo-Triassic?	S ₃ mylonitic fabric and
thrusting (C-SF thrust)		thrusting (C-SF thrust)
		S ₂ mylonitic fabric in quartzite
		and augen gneiss
		++++ Intrusion of protoliths of
		augen gneisses ++++ (~280-
		325 m.y. ago) Pennsylvanian
	pre-Pennsylvanian	S ₁ layering and F ₁ isoclines
		(Antler orogenic
		deformation?)

1981 RESEARCH

The Calaveras-Shoo Fly Thrust is separable into two major segments by the Standard pluton (Fig. 2). I have mapped the region north of the Standard pluton and show the data on Figure 3. The southern area is described below. I am convinced that the Shoo Fly exists as a continuous belt of polydeformed quartzitic gneiss, mica schist, diopsidic marble and augen gneiss as far south as 37°30'N where it is engulfed by granitic rocks of the Sierra Nevada batholith. The Shoo Fly, throughout the 1981study area (See 1981 RFP), is in ductile thrust contact with the Calaveras to the west and locally the thrust is overturned toward the west. Normally, it is a steep, east-dipping surface of structural discordance. Remarkably, the structural character of both the Calaveras and Shoo Fly are identical in the terranes both north and south of the Standard pluton. Table 1 can therefore be used for the southern area (study area) as well (Merguerian, 1982). What is more, the older fabrics in the Shoo Fly (S₁, S₂) are truncated at a high angle against the D₃, thrust zone. In addition, all of the various sub-units in the Shoo Fly are in fault contact with the Calaveras which suggests wholesale truncation of stratigraphic horizons in the Shoo Fly. The quadrangle mapping of 1981 is described below. I have tried to concentrate on discussion of geologic relations not shown on prevailing CDMG maps.

TUOLUMNE QUADRANGLE

The Calaveras-Shoo Fly Thrust is encountered southeast of the Standard pluton in sec. 10. It trends southeasterly to sec. 29 but is folded by N70°W-trending F₄ folds. In addition, the thrust is offset in secs. 24 and 19 by a high angle (EW, 50°N) brittle reverse fault that is more than likely a Cretaceous reactivation of a pre-existing S₄ surface. In the western part of the quadrangle, the thrust is folded and overturned, due to forceful intrusion of the Standard pluton. Reconnaissance mapping in the adjacent Standard quadrangle (Fig. 4) indicates the presence of Shoo Fly also overturned with respect to the Calaveras.

DUCKWALL MTN. QUADRANGLE

In this quadrangle the thrust enters and exits in the southwest quarter. It is strongly folded by E-W trending F₄ folds. In sec. 3 small isolated tectonic slivers of the Shoo Fly (garnet-staurolite-mica schist) are surrounded by highly sheared Calaveras chert. These occurrences are the result of imbrication of rock units during D₃ thrust development. The Shoo Fly underlies most, of the quadrangle although granitic rocks comprise the eastern portion. In secs. 24 and 25 a large plutonic mass of coarse-textured hornblendite, hornblende pyroxenite and subordinate gabbro are overlain by Tertiary pyroclastic mudflow breccia.

GROVELAND QUADRANGLE

All Calaveras rocks occur in the eastern part of the sheet. Jurassic phyllite occurs in the western half (reconnaissance) and is in ductile fault contact with the Calaveras to the east. Some refinement of Tertiary volcanic contacts are presented. Basically, little mapping was attempted on this sheet.

JAWBONE RIDGE QUADRANGLE

The Shoo Fly underlies the northeast portion of this sheet and the thrust is projected, with little control, into granitic rocks of the batholith in sec. 24 (Tuolumne River). While it is extremely unlikely that the thrust is not folded as shown, the region was inaccessible during the 1981 season. The placement of the thrust in the northern region is constrained by little outcrop data and geologic trends but F₄ digitations were not diagrammatically sketched although they most assuredly exist. The thrust re-emerges northeast of Pilot Ridge (sec. 32) where control is excellent and trends southeasterly to the edge of the sheet. A sliver of mylonitic ferruginous chert (Calaveras) was detected within the Shoo Fly in sec. 33. Early foliations in the Shoo Fly (sec. 33) are isoclinally folded, transposed and truncated at a high angle by the thrust.

LAKE ELEANOR NW QUADRANGLE

Mainly granite with local roof pendants and screens of Shoo Fly.

LAKE ELEANOR SW QUADRANGLE

The thrust enters this quadrangle in sec. 33 and trends southeasterly until it is crosscut by granodiorite of the Hazel Green pluton. The thrust surface is very complex in this region due to the superposition of E-W trending F_4 folds, NW-trending F_5 (Nevadan) folds, and a younger generation of NE-trending crenulate folds (F_6). Due to complexities during thrusting (i.e. - isoclinal and shear folding, transposition, development of blastomylonite, and tectonic imbrication of rock units, the thrust trace marks a form-line that separates $\geq 50\%$ Calaveras lithologies from $\geq 50\%$ Shoo Fly lithologies to the east. It is possible that in the vicinity of the thrust zone each outcrop is a fault-bounded sliver but I have opted to represent the contact, as a 50% form-line. Gabbroic plutons occur in the Shoo Fly in secs. 1, 6, 5, and 7.

LAKE ELEANOR SE QUADRANGLE

The Shoo Fly is limited to the southern quarter of this sheet, and the Calaveras does not exist. Most of the quadrangle is underlain by granitic rocks of the Sierra Nevada batholith. See Tobisch (1960ms) for more data.

EL PORTAL 15' QUADRANGLE

The El Portal quadrangle is a compilation of mapping done on the Kinsley, El Portal and Buckingham Mountain quadrangles during 1981. In addition, Dr. Richard A. Schweickert has included data from parts of the Feliciana Mountain quadrangle to complete the 15' sheet.

The thrust emerges from the Hazel Green pluton in sec. 27 and trends southeasterly past the Merced River. Control is good in this segment and post-thrust folding (F_4 ?) can be seen south of sec. 35. The folds are tentatively identified as F_4 folds that have been rotated by local intrusives. At the Merced River, imbrication of the Shoo Fly is obvious as well as intense mylonitization of both the Shoo Fly and Calaveras complexes. The thrust is dashed to the south

of the Merced River due to inaccessibility in 1981 but the thrust is accurately mapped in the extreme southeast part of the quadrangle (secs. 10 and 15) where it is truncated by the batholith.

The Calaveras occupies the central part of the sheet and is inductile fault contact with Jurassic phyllite and greenstone to the west. Complications exist on Buckingham Mountain as rocks there are highly flattened and are lithically identical to the Shoo Fly occurring to the east. This region needs to be re-examined and will be a focal point for a 1982 RFP to be submitted shortly.

In summary, the extent of the Shoo Fly and the Calaveras-Shoo Fly Thrust has now been mapped as far south as Lat. 37°30'N and shows an identical structural signature as found north of the Standard pluton. The existence of the Shoo Fly in this region poses questions concerning the extent of the Shoo Fly in roof pendants to the east, southeast, and south of the 1981 study area.

Respectfully submitted, Charles Merguerian 22 December 1981

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Note: Map sheets available from the California Division of Mines and Geology or http://www.dukelabs.com. Map explanations are below.

Filename: CM1981.doc

LITHOLOGIC UNITS

Pleistocene Non-Marine Deposits

Ta Tertiary Non-Marine Deposits

Pliocene Mehrten Formation - pyroclastic rocks

Granitic to granodioritic rocks of the Sierra Nevada batholith

U Ultramafic plutonic rocks

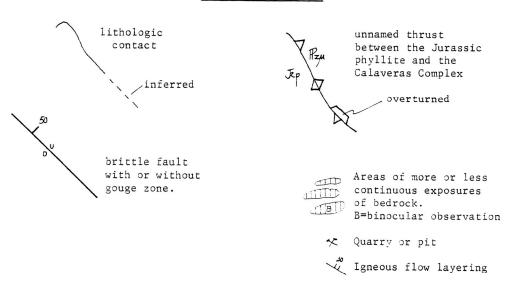
96 Gabbroic rocks

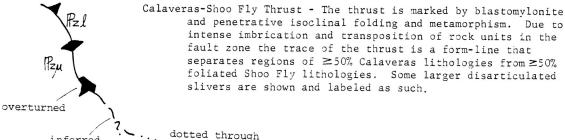
Æρ Jurassic phyllite Æθ Jurassic greenstone

Upper Paleozoic Calaveras Complex - argillite, chert-argillite, Pzu rhythmically-bedded and massive chert, marble (ls), talc-schist, rare sandstone layers, post-tectonic mafic dikes. The Calaveras Complex is in ductile fault contact with both the Jurassic rocks to the west and with the Shoo Fly Complex to the east.

Lower Paleozoic Shoo Fly Complex - dominantly mylonitic quartzite Pzl and micaceous quartzofeldspathic gneiss and mica-quartz gneiss and schist $^{+}$ garnet and graphite. In addition diopsidic calcsilicate and marble+ biotite, graphitic marble and schist. In Tuolumne County, the Shoo Fly also contains numerous bodies of Paleozoic intrusives that now occur as augen gneiss (ag) and many post-tectonic mafic dikes.

LITHOLOGIC CONTACTS





intrusive rocks

inferréd

STRUCTURAL SYMBOLS

The symbols are defined for each lithotectonic belt. Symbols can be mixed. The point of intersection is the observation point.



Jurassic phyllite and greenstone belt

S₁ - Dominant phyllitic cleavage and lithologic layering

Upper Paleozoic Calaveras Complex

- S_0 Bedding S_1 Flattening foliation
- F_1 , L_1 Fold axis/stretching lineation S_2 Slip cleavage with recrystallized biotite
- \mathbf{F}_2 , \mathbf{L}_2 Fold axis/crenulation or intersection lineation

Lower Paleozoic Shoo Fly Complex

- S_0 compositional layering (bedding) proves rocks are metasedimentary but of little structural help due to extensive shearing and transposition.
- S_1 or S_2 Pre-thrust event schistosity or gneissosity with two superposed amphibolite facies folding episodes (F1, F_2). These are often mylonitic. S_1 is rare and is locally noted on the maps.
- F_1 , F_2 , L_1 , L_2 fold axis / stretching lineation S₃ - blastomylonitic amphibolite facies foliation formed during development of the Calaveras - Shoo Fly Thrust. Within 1 km of the ductile fault, shearing, boudinage, transposition and metamorphic overprinting of the older fabrics $(S_1 \text{ and } S_2)$ is most complete. Away from the thrust zone the S_3 foliation is a penetrative mica foliation with shearing along $S_{\mathfrak{J}}$ axial surfaces. On the maps the S_3 foliation symbol was \underline{only} plotted where field relations were unequivocal. The older fabrics are best developed away from the thrust within the main body of the Shoo Fly.
- \mathbf{F}_3 , \mathbf{L}_3 Fold axis / stretching lineation S₄ - slip cleavage with recrystallized biotite and a local black residue cleavage consisting of biotite + opaque minerals. F_4 , L_4 - Fold axis / crenulation or intersection lineation S_5 , S_6 , S_7 - younger slip cleavages
 - $\rm S_{5}$ N32W, 78 $^{\rm O}$ NE (Nevadan cleavage with crenulate to open folds (F₅)
 - S₆ N30E, 90° (Late Nevadan cleavage with crenulate to open folds (F₆)
 - S, N70W EW, 90° (Cretaceous high angle rock cleavage and local reverse faulting with quartz veining and mineralization). Local open folds (F_7) observed, however, F5, 6, or 7 folds are generally not plotted.

Geology mapped in July and August, 1981 by Charles Merguerian